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Simultaneous high-pressure and high-temperature volume measurements of ice VII and its thermal equation of state

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We conducted simultaneous high-pressure (P) and high-temperature (T) unit-cell volume (V) measurements of H2O ice VII in an externally heated diamond anvil cell using in situ synchrotron x-ray diffraction technique with an angle-dispersive system. Isothermal unit-cell volume of ice VII was collected at P = 33–50 GPa with T = 873 K. Nonisothermal experiments at T = 430–740 K and P = 19–37 GPa were also conducted to further constrain P-V-T properties of ice VII. Using the existing 300 K compression data for an equation of state (EoS) at the reference temperature, a thermal EoS of ice VII was constructed from the collected high-P-T data. The melting temperature of ice VII was computed up to P = 40 GPa by thermodynamic calculations with the newly constructed EoS of ice VII. The resulting melting temperatures are higher than that of previous calculations since the new EoS gives a smaller molar volume of ice VII. In contrast, the calculated melting curves are consistent with recent experimental estimates.

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I. INTRODUCTION

Physical properties of H2O at high-pressure (P) and high-temperature (T) have been under deep interest in condensed matter physics and chemistry, and planetary sciences. Equation of state (EoS) of solid H2O plays an important role in understanding hydrogen bond behavior under extreme P-T conditions. Melting curve of H2O is the primary information for modeling the deep interiors of icy planets such as Neptune and Uranus,1 and for discussing the possible existence of H2O solid phase in Earth’s deep mantle that may have been delivered by cold subducting slabs.2

Ice VII is one of the 15 known solid phases of H2O, and it appears above P = 2.2 GPa at room temperature along compression after the liquid to ice VI transformation at P = 0.9 GPa. Ice VII takes a body-centered-cubic (bcc) oxygen sublattice and this simple structure is preserved to a very high pressure of 170 GPa.3 The experimental evidence for the oxygen sublattice is preserved through the phase transitions which are related to the hydrogen-bond symmetrization, the EoS of ice must be evaluated separately for each phase. More importantly, their study suggested that in order to observe such a change in compressibility, the high-resolution XRD technique should be used. Recently a study by the quantum molecular-dynamics simulations17 reproduced the 300 K data set of Sugimura et al.16 and therefore confirmed the validity of their experimental observations.

In contrast to 300 K experiments, high-temperature experimental volumetric data for high-pressure ice are available only from two energy-dispersive XRD works. Fei et al.18 collected the high-P-T volume data of ice at P = 6–20 GPa and T = 300–650 K, and the same group, Frank et al.19 later extended the maximum pressure and temperature up to 60 GPa and 850 K. However, as mentioned above the use of the angle-dispersive XRD technique is required to observe a precise P-V relation of high-pressure ice. In addition, it is highly plausible that Frank et al.19 included the data of dynamically disordered ice VII, which is significantly compressible compared with ice VII,16 when constructing the P-V-T EoS of ice VII.

The aim of the present study is to construct the P-V-T EoS of ice VII from the volumetric data collected only from
II. EXPERIMENTAL PROCEDURES

High-$P$-$T$ conditions were generated in the Bayreuth-type externally heated diamond anvil cell (EHDAC). For the heater, we used a cylindrical furnace where AlChrom-O wire (0.25 mm in diameter) was wound. Heating was performed by supplying electricity to the AlChrom-O wire from a dc power supply up to 110 V to reach 873 K. Temperature was measured by Alumel-Chromel ($K$-type) thermocouple, which was placed very close (about 2 mm) to the sample. Due to a large heater size compared with the sample, uniform heating across the sample chamber was achieved. The temperature fluctuation during the heating was very small across the sample chamber was achieved. The temperature uncertainty was less than 5 K. The entire temperature uncertainty was less than 5 K. The Bayreuth type EHDAC works at BL10XU, SPring-8.

Deionized liquid H$_2$O was loaded into a hole drilled in a preindented rhenum gasket with fine powder of gold (Au) or platinum (Pt) used as an internal pressure calibrant. With flat 300 $\mu$m or beveled 200 $\mu$m culet diamond anvils, the sample was first compressed to several tens of gigapascals at room temperature. Temperature was increased by referring to the previously reported melting curves of H$_2$O ice, although some measurements were conducted at temperatures above the several melting curves among those listed above. Details will be discussed later in Sec. IV.

Angle-dispersive XRD spectra were collected on an imaging plate (Rigaku-R-AXIS IV) at BL10XU of SPring-8. Monochromatic incident x-ray beam with a wavelength of 0.411–0.415 Å was collimated to 15–20 $\mu$m in diameter. The exposure time was 2 min. Two-dimensional (2D) XRD images were integrated as a function of $2\theta$ angle using the fit-2D program to obtain conventional one-dimensional diffraction patterns.

III. RESULTS

A. Experimental results

Total of five separate sets of high-$P$-$T$ in situ XRD experiments were conducted. The $P$-$T$ conditions for volume measurements are shown in Fig. 1. A typical XRD spectrum of the sample is illustrated in Fig. 2. The unit-cell volumes of ice were determined from (110) and (200) peaks of bcc structure for all the XRD data. The splitting of the diffraction peak of ice was not observed at any $P$-$T$ condition, which may contradict a recent experimental study at $T=300$ K by Somayazulu et al. This discrepancy will be discussed in Sec. IV.

The pressures were determined from the unit-cell volume of Au or Pt calculated from its (111), (200), and (220) diffraction lines. Sugimura et al. used Pt as a pressure calibrant in 300 K experiments and their 300 K EoS was employed as a reference in this study. However, from a laser-heated DAC experiment, it is known that Pt may react with hydrogen to form a platinum hydride, PtH$_X$ at $T=1350–1500$ K. Hence, although the present experiments do not reach such high-$T$ conditions, we employed Au which is expected to be less chemically reactive than Pt, as the main pressure calibrant and Pt as the secondary pressure calibrant throughout the experiments. In order to sustain the consistency between different calibrants, we employed Au’s self-consistent pressure scales for Au and Pt which should give the identical pressure. As shown in Fig. 2, we did not observe any XRD peaks that signify the formation of PH$_X$ at all experimental $P$-$T$ conditions.

The experimental conditions and the unit-cell volumes of H$_2$O ice are summarized in Table I. Isothermal $P$-$V$ data at


TABLE I. Experimental conditions and the volumes of H$_2$O ice VII.

<table>
<thead>
<tr>
<th>Data</th>
<th>T (K)</th>
<th>P (GPa)</th>
<th>$V_{he VII}$ (cm$^3$/mol)</th>
<th>$a_{Ice VII}$ (Å)</th>
<th>$a_{Au}$ (Å)</th>
<th>$a_{Pt}$ (Å)</th>
<th>$S_t$</th>
</tr>
</thead>
<tbody>
<tr>
<td>01-T16</td>
<td>431</td>
<td>35.60 ± 0.07</td>
<td>7.3378 ± 0.0002</td>
<td>2.8996 ± 0.0000</td>
<td>3.8939 ± 0.0003</td>
<td>0.0003</td>
<td></td>
</tr>
<tr>
<td>02-T15</td>
<td>449</td>
<td>37.60 ± 0.30</td>
<td>7.2315 ± 0.0243</td>
<td>2.8855 ± 0.0032</td>
<td>3.8871 ± 0.0011</td>
<td>0.0024</td>
<td></td>
</tr>
<tr>
<td>03-T14</td>
<td>467</td>
<td>37.19 ± 0.07</td>
<td>7.2906 ± 0.0122</td>
<td>2.8933 ± 0.0016</td>
<td>3.8890 ± 0.0002</td>
<td>0.0006</td>
<td></td>
</tr>
<tr>
<td>04-T15</td>
<td>565</td>
<td>28.42 ± 0.11</td>
<td>7.7935 ± 0.0288</td>
<td>2.9584 ± 0.0036</td>
<td>3.9264 ± 0.0004</td>
<td>0.0015</td>
<td></td>
</tr>
<tr>
<td>05-T16</td>
<td>596</td>
<td>34.03 ± 0.14</td>
<td>7.4923 ± 0.0042</td>
<td>2.9198 ± 0.0005</td>
<td>3.9046 ± 0.0005</td>
<td>0.0008</td>
<td></td>
</tr>
<tr>
<td>06-T15</td>
<td>701</td>
<td>27.29 ± 0.16</td>
<td>7.9492 ± 0.0173</td>
<td>2.9780 ± 0.0022</td>
<td>3.9356 ± 0.0007</td>
<td>0.0024</td>
<td></td>
</tr>
<tr>
<td>07-T15</td>
<td>711</td>
<td>29.72 ± 0.14</td>
<td>7.7995 ± 0.0130</td>
<td>2.9591 ± 0.0016</td>
<td>3.9255 ± 0.0006</td>
<td>0.0020</td>
<td></td>
</tr>
<tr>
<td>08-T16</td>
<td>730</td>
<td>19.36 ± 0.13</td>
<td>8.6262 ± 0.0091</td>
<td>3.0602 ± 0.0011</td>
<td>3.9744 ± 0.0007</td>
<td>0.0022</td>
<td></td>
</tr>
<tr>
<td>09-T16</td>
<td>740</td>
<td>18.88 ± 0.05</td>
<td>8.6910 ± 0.0068</td>
<td>3.0678 ± 0.0008</td>
<td>3.9773 ± 0.0003</td>
<td>0.0005</td>
<td></td>
</tr>
<tr>
<td>10-T16</td>
<td>740</td>
<td>22.04 ± 0.11</td>
<td>8.3774 ± 0.0196</td>
<td>3.0305 ± 0.0024</td>
<td>3.9614 ± 0.0006</td>
<td>0.0018</td>
<td></td>
</tr>
<tr>
<td>11-T16</td>
<td>740</td>
<td>24.33 ± 0.12</td>
<td>8.2695 ± 0.0107</td>
<td>3.0174 ± 0.0013</td>
<td>3.9546 ± 0.0006</td>
<td>0.0017</td>
<td></td>
</tr>
<tr>
<td>12-T16</td>
<td>741</td>
<td>20.06 ± 0.10</td>
<td>8.5665 ± 0.0026</td>
<td>3.0531 ± 0.0003</td>
<td>3.9712 ± 0.0005</td>
<td>0.0010</td>
<td></td>
</tr>
<tr>
<td>13-T09</td>
<td>867</td>
<td>50.25 ± 0.30</td>
<td>8.6932 ± 0.0150</td>
<td>2.8398 ± 0.0021</td>
<td>3.7667 ± 0.0008</td>
<td>0.0003</td>
<td></td>
</tr>
<tr>
<td>14-T09</td>
<td>868</td>
<td>49.43 ± 0.12</td>
<td>6.9346 ± 0.0119</td>
<td>2.8455 ± 0.0016</td>
<td>3.7688 ± 0.0003</td>
<td>0.0001</td>
<td></td>
</tr>
<tr>
<td>15-T16</td>
<td>870</td>
<td>36.50 ± 0.17</td>
<td>7.4511 ± 0.0176</td>
<td>2.9144 ± 0.0023</td>
<td>3.9030 ± 0.0006</td>
<td>0.0021</td>
<td></td>
</tr>
<tr>
<td>16-T16</td>
<td>870</td>
<td>38.72 ± 0.26</td>
<td>7.3534 ± 0.0206</td>
<td>2.9016 ± 0.0027</td>
<td>3.8946 ± 0.0010</td>
<td>0.0032</td>
<td></td>
</tr>
<tr>
<td>17-T16</td>
<td>870</td>
<td>44.24 ± 0.26</td>
<td>7.1063 ± 0.0298</td>
<td>2.8687 ± 0.0040</td>
<td>3.8748 ± 0.0009</td>
<td>0.0032</td>
<td></td>
</tr>
<tr>
<td>18-T09</td>
<td>870</td>
<td>47.62 ± 0.32</td>
<td>6.9924 ± 0.0130</td>
<td>2.8533 ± 0.0018</td>
<td>3.7735 ± 0.0008</td>
<td>−0.0007</td>
<td></td>
</tr>
<tr>
<td>19-T09</td>
<td>870</td>
<td>49.12 ± 0.43</td>
<td>6.9467 ± 0.0138</td>
<td>2.8471 ± 0.0019</td>
<td>3.7696 ± 0.0011</td>
<td>−0.0020</td>
<td></td>
</tr>
<tr>
<td>20-T16</td>
<td>871</td>
<td>46.43 ± 0.27</td>
<td>7.0239 ± 0.0286</td>
<td>2.8576 ± 0.0039</td>
<td>3.8675 ± 0.0009</td>
<td>0.0031</td>
<td></td>
</tr>
<tr>
<td>21-T16</td>
<td>872</td>
<td>33.87 ± 0.16</td>
<td>7.5980 ± 0.0265</td>
<td>2.9334 ± 0.0034</td>
<td>3.9134 ± 0.0006</td>
<td>0.0022</td>
<td></td>
</tr>
<tr>
<td>22-T16</td>
<td>873</td>
<td>32.78 ± 0.24</td>
<td>7.6637 ± 0.0092</td>
<td>2.9419 ± 0.0012</td>
<td>3.9178 ± 0.0010</td>
<td>0.0027</td>
<td></td>
</tr>
<tr>
<td>23-T19</td>
<td>875</td>
<td>48.11 ± 0.19</td>
<td>6.9734 ± 0.0319</td>
<td>2.8508 ± 0.0043</td>
<td>3.8620 ± 0.0006</td>
<td>0.0019</td>
<td></td>
</tr>
<tr>
<td>24-T19</td>
<td>875</td>
<td>49.36 ± 0.40</td>
<td>6.9303 ± 0.0180</td>
<td>2.8449 ± 0.0025</td>
<td>3.8580 ± 0.0013</td>
<td>0.0045</td>
<td></td>
</tr>
<tr>
<td>25-T19</td>
<td>877</td>
<td>45.61 ± 0.11</td>
<td>7.0599 ± 0.0109</td>
<td>2.8625 ± 0.0015</td>
<td>3.8703 ± 0.0004</td>
<td>0.0012</td>
<td></td>
</tr>
<tr>
<td>26-T19</td>
<td>881</td>
<td>45.78 ± 0.11</td>
<td>7.0521 ± 0.0032</td>
<td>2.8614 ± 0.0004</td>
<td>3.8699 ± 0.0004</td>
<td>0.0006</td>
<td></td>
</tr>
</tbody>
</table>

$T$ = 867–880 K (−873 K) were collected at pressure from 33 to 50 GPa. In order to further constrain the $P$-$V$-$T$ properties of ice VII, additional volume data were also collected at $T = 431$–741 K and $P = 19$–38 GPa. Figure 3 shows the collected volumes of H$_2$O ice plotted against the pressures, together with the isothermal compression data of ice VII at $T = 300$ K. The isothermal $P$-$V$ data at $T = 873$ K run almost parallel to the 300 K isotherm of ice VII, suggesting the data were collected well within the stability field of ice VII. Namely, an anomalous volume reduction, the plausible signature of the phase transition of ice VII to dynamically disordered ice VII, did not occur at the pressure range of the present experiments. Hence, it can be inferred that ice VII is stable at $T = 873$ K at least to $P = 50$ GPa.

In the high-pressure experiments using a DAC, the hydrostatic state of the sample must be somewhat evaluated, because it is a significant issue concerning the compression behavior of solid materials, especially since pressure medium was not used in the present experiments. We followed the method proposed by Shim et al. who used $S_t$ value as an indicator of the magnitude of the uniaxial stress component in a cubic sample; $S$ is the elastic anisotropy and $t$ is the uniaxial stress component (see Shim et al. for details). Here, $S_t$ value of each experimental data was calculated from (111), (200), and (220) peaks of Au or Pt mixed with the sample. The estimated $S_t$ for each data point is listed in Table I. In our $P$-$V$-$T$ data set, the $S_t$ value ranges from −0.0020 to −0.0003.

![FIG. 3. Molar volumes of H$_2$O ice VII. Isotherms are calculated at $T = 300$, 473, 673, and 873 K from EqS with parameters listed in Table II. The circles are the high-temperature data collected in this study and the triangles are the 300 K experimental data of ice VII from Sugimura et al. (Ref. 16).](image-url)
TABLE II. Thermoelastic parameters of H$_2$O ice VII.

<table>
<thead>
<tr>
<th></th>
<th>$V_0$ (cm$^3$/mol)</th>
<th>$K_0$ (GPa)</th>
<th>$K'$</th>
<th>$\alpha_0$ (10$^{-5}$/K)</th>
<th>$\delta_T$</th>
</tr>
</thead>
<tbody>
<tr>
<td>This study</td>
<td>14.52$^a$</td>
<td>5.02</td>
<td>7.51</td>
<td>150 ± 19</td>
<td>5.1 ± 0.2</td>
</tr>
<tr>
<td>Fei et al.$^b$</td>
<td>12.3 ± 0.2</td>
<td>23.9 ± 0.7</td>
<td>4.2  ± 0.5</td>
<td>6.0</td>
<td></td>
</tr>
<tr>
<td>Frank et al.$^c$</td>
<td>12.4 ± 0.1</td>
<td>21.1 ± 1.3</td>
<td>4.4 ± 0.1</td>
<td>4.8</td>
<td></td>
</tr>
</tbody>
</table>

$^a$From Loubye et al. (Ref. 3).
$^b$Reference 18.
$^c$Reference 19.

0.0045. If $S$ is available, one can calculate the deviatoric stress from $S$. However, since $S$ varies with pressure (e.g., Duffy et al.$^{35}$) and its pressure dependence for Au and Pt is unknown, we cannot directly estimate the deviatoric stress in the sample. For reference, Shim et al.$^{36}$ reported the $[S]$ value of CaSiO$_3$ perovskite at $P$=18–96 GPa and $T$=1238–2419 K ranged from $-0.005$ to 0.005. In addition, Dubrovinsky et al.$^{37}$ reported that deviatoric stress in MgO is as small as 0.25 GPa at $T$=850 K. Therefore, the stress state in the present EHDAC was expected to be minimal.

B. P-V-T equation of state of ice VII

The P-V-T EoS of ice VII was constructed from the high-temperature data collected in this study. The existing EoS at $T$=300 K (Ref. 16) which was obtained by the same experimental technique as in this study was used as the reference temperature EoS. The molar volumes of ice VII at $T$=300 K are described with the Vinet EoS

$$P = 3K_0\alpha_0 x^2(1-x)\exp\left[\frac{3}{2}(K'-1)(1-x)\right],$$

where $x=(V/V_0)^{1/3}$, and $V_0$, $K_0$, and $K'$ are the molar volume, isothermal bulk modulus, and its pressure derivative, respectively, at ambient condition. Here, $V_0$=14.52 cm$^3$/mol, $K_0$=5.02 GPa, and $K'$=7.51 for ice VII from Sugimura et al.$^{16}$ are adopted.

In order to extend EoS to high temperatures, we express the thermal expansivity as a function of pressure in the framework of the Anderson-Grüneisen expression,$^{38}$ described as

$$\frac{\alpha}{\alpha_0} = (V/V_0)^{\delta_T},$$

where $\alpha$ and $\alpha_0$ are the thermal expansivity at a pressure of interest and ambient pressure, respectively. The parameter $\delta_T$ is so-called Anderson-Grüneisen parameter. A least-square fitting of our high-pressure and high-temperature data yielded $\alpha_0=150 \pm 19 \times 10^{-5}$/K and $\delta_T=5.1 \pm 0.2$. The obtained parameters are summarized in Table II. In Fig. 4, the calculated isotherms at every 100 K of ice VII up to $T$=900 K are illustrated together with collected high-temperature volume data, in terms of the difference from those of room temperature at corresponding pressure, $V(P,T)-V(P,300 \text{ K})$. Figure 4 emphasizes that the experimental data are well reproduced by the P-V-T EoS with the obtained parameters.

C. Melting curve of ice VII

Melting of ice VII is defined by the equilibrium of the reaction of H$_2$O (ice VII)=H$_2$O (fluid). On the equilibrium $P$-$T$ conditions, the Gibbs free energy difference between the two phases is zero, namely,

$$\Delta G_{(\text{Ice VII-Water})} = 0 = \Delta H_{T} + \int_{T_0}^{T} \Delta C_{P,T} dT
- T \left[ \frac{\Delta S_{T}}{T} + \int_{T_0}^{T} \frac{\Delta C_{P,T}}{T} dT \right] + \int_{P_0}^{P} \Delta V_{T} dP,$$

where $\Delta G_{T}$ is a Gibbs free energy change in the reaction, $\Delta H_{T}$, $\Delta S_{T}$, $\Delta C_{P,T}$, and $\Delta V_{T}$, are the changes in the reaction in enthalpy, entropy, heat capacity, and molar volume, respectively. The reference $P$-$T$ condition is 1 bar and 300 K. The superscript O on each parameter denotes 1 bar. The $\Delta H_{T}$ and $\Delta S_{T}$ are calculated from the difference between each phase in the standard enthalpy of formation from the elements ($\Delta H_{O}^{\text{f}}$) and standard entropy ($S_{O}^{0}$), both of which are at the reference $P$-$T$ conditions. The Gibbs free energy of fluid H$_2$O at $P$=1 bar and $T$ of interest is given by Robie et al.$^{39}$

FIG. 4. Molar volumes of H$_2$O ice VII relative to the volumes at 300 K. Isotherms (solid lines) are calculated from the EoS in this study in every 100 K up to 900 K. The horizontal line represents the reference isotherm that is of 300 K (Ref. 16). The dashed line represents the calculated 873 K isotherm, which well reproduces the 873 K experimental data (closed circles).
Using their $P$-$V$-$T$ EoS for ice VII, Fei et al.\textsuperscript{18} obtained $\Delta_H^0$, $S^0$, and $C_p^0$ for ice VII by fitting previously published experimental data\textsuperscript{20} up to $P$=5 GPa. From the determined parameters, they calculated the melting temperatures of ice VII at pressures from 5 to 20 GPa, which are fairly consistent with the result of their melting experiments. We revisited the thermodynamic calculation of the melting of ice VII by replacing the $P$-$V$-$T$ EoS for ice VII with that in this study. The present purpose is to evaluate the effect of $P$-$V$-$T$ EoS for ice VII on the calculated melting temperature. For $\Delta_H^0$, $S^0$, and $C_p^0$ of fluid H\textsubscript{2}O, we adopted the data in Robie et al.\textsuperscript{39} same as in Fei et al.\textsuperscript{18} The heat capacity for ice VII was also from Fei et al.\textsuperscript{18} The specific values of each parameter are summarized in Table III. For the calculation of fluid fugacity, we used EoS proposed by Halbach and Chatterjee\textsuperscript{40} and Belonoshko and Saxena\textsuperscript{41} which were used in Fei et al.\textsuperscript{18} In addition, we tested another EoS by Brodholt and Wood.\textsuperscript{42} We derived $\Delta_H^0$ and $S^0$ for ice VII to reproduce the experimental data of melting temperature up to $P$=5 GPa.\textsuperscript{20} Note that the set of $\Delta_H^0$ and $S^0$ was obtained for each of three calculations with different EoS of fluid H\textsubscript{2}O. Figure 5 shows the calculated melting temperatures of ice VII. Comparison of the curves between this study and Fei et al.\textsuperscript{18} shows that the melting temperature in this study is slightly lower at pressures below 13 GPa (with fluid EoS by Halbach and Chatterjee\textsuperscript{40}) and 11 GPa (with fluid EoS by Belonoshko and Saxena\textsuperscript{41}). However, beyond those pressures, our melting curve gives higher temperatures.

### IV. DISCUSSIONS

#### A. $P$-$V$-$T$ equation of state of ice VII

Newly determined $P$-$V$-$T$ EoS of ice VII in this study is compared with that of previous studies\textsuperscript{18,19} in Table II and Fig. 6. Our thermal expansivity at $P$=1 bar ($\alpha_0$) is very large compared with that of the previous estimates\textsuperscript{18,19} (Table II). This apparent discrepancy comes from the difference in the formulation of EoS at $T$=300 K. Since ice VII is not quenchable to $P$=1 bar, the 1 bar parameters of $V_0$, $K_0$, and $\alpha_0$ are obtained as the fitted parameters depending on the formulation of EoS. Previous works\textsuperscript{18,19} adopted a Birch-Murnaghan EoS. Small $K_0$ in this study is due to the use of the Vinet EoS, which resulted in relatively large $\alpha_0$. Thus, the difference in the value of $\alpha_0$ between the present study and the previous works is not essential.

Fei et al.\textsuperscript{18} evaluated the $P$-$V$-$T$ EoS of ice VII from the \textit{in situ} XRD experiments in an EHDAC at $P$=6–20 GPa and $T$=300–700 K. After 10 years, the same group\textsuperscript{19} extended the $P$-$T$ range of the experiment to 65 GPa and 850 K. The difference from our EoS reaches to 5.2% at $T$=900 K.\textsuperscript{18} As mentioned above, both of the previous studies used an energy-dispersive XRD technique whose resolution is lower than the angle-dispersive technique taken in this study. Sugimura et al.\textsuperscript{16} showed that the angle-dispersive XRD technique allowed one to detect the phase transitions in ice which had not been observed in the energy-dispersive techniques. In addition to the difference in XRD system, we discuss other possible sources for the discrepancies below.

![Melting curve of H\textsubscript{2}O ice VII](image-url)

**FIG. 5.** (Color online) Melting curve of H\textsubscript{2}O ice VII. The red curves are the melting lines of ice VII calculated in this study with the EoS of fluid H\textsubscript{2}O of Brodholt and Wood (Ref. 42) (upper solid line) and Belonoshko and Saxena (Ref. 41) (lower solid line), and Halbach and Chatterjee (Ref. 40) (dashed line). The two solid red curves are taken as the upper and lower bounds in this study. The black solid and dashed lines are the result of Fei et al.’s (Ref. 18) calculations with the EoS of fluid H\textsubscript{2}O from Belonoshko and Saxena (Ref. 41) and Halbach and Chatterjee (Ref. 40), respectively. The black thin dotted lines are the melting curves of ice VII proposed by previous experimental studies (Refs. 18–26) shown by letters as described in Fig. 1.
FIG. 6. Comparison of the EoS of this study (0%) with that of the previous studies (Refs. 18 and 19).

For the comparison with Fei et al.'s18 volumes of ice VII calculated from their EoS are larger than those calculated from our EoS by 0.0–2.2 %, 0.2–1.8 %, and 2.0–5.2 %, at $T=300$, 600, and 900 K, respectively, in the pressure range shown in Fig. 6.

In order to compare the high-$P$-$T$ experimental results, attention should be paid in pressure scale. Fei et al.'s18 calculated their pressures using the Anderson et al.'s44 Au scale while the present experiments determined the pressures from Fei et al.'s Au or Pt scale. It has been pointed out that Anderson et al.'s44 Au scale underestimates the pressure compared with other Au scales, e.g., 2.4 GPa lower than Fei et al.'s Au scale at $P=50$ GPa and $T=873$ K. However, Fig. 6 shows that the volumes of ice VII by Fei et al.'s18 are larger than those by this study, namely, they overestimated the pressures. Thus, the discrepancy in volume of ice VII cannot be explained by the pressure scale difference.

Another possible source for the discrepancy is the nonhydrostatic environments in each experiment. In general, the nonhydrostatic condition causes an overestimation of the sample volume as shown previously, e.g., MgO.33 Our $P$-$V$-$T$ EoS of ice VII is constructed from the 300 K EoS from Sugimura et al.'s16 and the high-temperature volume data collected in the present experiments. Sugimura et al.'s16 evaluated the nonhydrostatic stresses on the sample via difference in lattice parameters of ice, determined separately from (110) and (200) diffraction lines. The difference does not exceed ~0.4 % even at $P=55$ GPa, which is satisfactorily small concerning the nonhydrostatic stress on the sample at $T=300$ K. For the present high-temperature experiments, the nonhydrostatic stress should be small due to the high temperature.

According to Dubrovinsky et al.,37 the nonhydrostatic stress can be reduced by increasing temperature and it almost vanishes at $T\geq 800$ K in a DAC at $\sim 30$ GPa. In this study, about 50 % of our $P$-$V$-$T$ data were collected at temperature well above 800 K (Table 1) while Fei et al.'s18 experiments were at $T\leq 650$ K, suggesting large portion of their $P$-$V$-$T$ data were collected under highly nonhydrostatic conditions. Therefore, it is quite plausible that larger volumes of ice VII in Fei et al.'s18 were caused by the nonhydrostatic stress on their sample.

The EoS of Frank et al.'s19 gives volume of ice VII in better agreement with our data than Fei et al.'s18. As in the case of Fei et al.'s18, the nonhydrostatic stress may account for the large volume of Frank et al.'s19 since only about 9 % of Frank et al.'s19 high-temperature data set are from $T=800$ K. Since all the Fei et al.'s18 experiments were made at $T\leq 650$ K, the data set of Frank et al.'s19 should have been taken under better hydrostatic conditions than that of Fei et al.'s18. This is evidenced by the fact that EoS of Frank et al.'s19 gives closer volume to ours than what Fei et al.'s18 EoS gives. The hydrostaticity of the sample in our experiment is superior to the others.

At high pressures, the crossover occurs and the volumes by Frank et al.'s19 eventually become the smallest before the pressure exceeds 48 GPa. The crossover should be accounted for the inclusion of the volume data from pressures above 40 GPa at room temperature or higher pressures at high-temperature conditions. At those high pressures, ice VII might have transformed into dynamically disordered ice VII,16 whose volume is smaller than ice VII.

Finally, we should note a recent work by Somayazulu et al.,15 which proposed a ferroelastic phase transition in ice VII at $P=14$ GPa and $T=300$ K from the fact that they observed the peak broadening in XRD patterns of ice. Since present (high-$T$) and previous (300 K) (Ref. 16) studies did not collect the data at $P<14$ GPa, the two data set cannot directly address the change in the peak width at $P=14$ GPa, although both studies did not observe any sign of the peak broadening of ice VII at any $P$-$T$ condition. Note that the peak broadening at $P=14$ GPa has never been observed in the other XRD studies as well, even with an angle-dispersive technique.7,23,50 The peak broadening in Somayazulu et al.'s15 could be a result of large uniaxial stress in a DAC experiment although they conducted laser annealing. The laser heating technique should have produced a large temperature gradient particularly in the compressional direction because they did not put a thermal insulator between the sample and diamond. Therefore, the stress in the sample might have been only partially released upon laser annealing. Moreover, Somayazulu et al.'s15 used an x-ray beam with a very large diameter in order to probe the entire sample chamber so that the pressure gradient in the radial direction was also sampled. Those pressure gradients must have caused a
peak broadening. In contrast, our heating method is the external heating with a very large heater and the diameter of the x-ray beam is 15–25% of that of the sample chamber. Our technique provides better hydrostatic environment in the probed area than Somayazulu et al.'s.

B. Melting curve of ice VII

The melting temperature of ice VII at high pressures has been a matter of debate. As listed above, the numbers of melting curve with different melting criteria have been proposed. Among those proposed criteria, the disappearance of XRD peaks of solid phase ice VII was criticized due to the possibility of the grain growth. Near the melting temperatures, the ice grains could rapidly grow and possibly be oriented in a direction such that the XRD appears to be absent. Fig. 1 shows that we observed solid ice VII even in the fluid stability field in Fei et al. and Frank et al. They judged melting from presence/absence of XRD peaks. Thus, Fei et al. and Frank et al. underestimated the stability of solid ice VII. Indeed, another melting curve determined by the same XRD method is located at higher temperature than Fei et al. and Frank et al. These discrepancies from the same method are likely due to the grain growth of ice.

The calculated melting temperature in this study is higher than the previously calculated one above P=13 GPa (Fig. 5), which comes from the difference in the P-V-T EoS for ice VII. The present experimental P-T conditions are within the calculated stability field of ice VII, supporting the above discussion that the P-V-T EoS for ice VII constructed in this study is more reliable than in Fei et al.

The thermodynamic calculation for the melting of ice depends also on the P-V-T EoS for fluid H2O. Among the numbers of proposed EoS for fluid H2O, the EoS by Belonoshko and Saxena has been widely used due to its wide applicable P-T range (up to 100 GPa and 4000 K). Here we tested another EoS by Brodholt and Wood which is also applicable in a wide P-T range (up to 30 GPa and 2000 K). Komabayashi et al. argued that the fluid EoS of Brodholt and Wood reproduced experimental data of a dehydration reaction of a hydrous phase up to P=8 GPa and T=1100 K and therefore, concluded this EoS was reliable. The calculated melting temperatures of ice VII differ by 300 K at P=30 GPa between two fluid EoS (Refs. 41 and 42) (Fig. 5). We adopted this difference as the error bar in our calculations. The melting curves by Datchi et al. and Lin et al. are located in this uncertainty range. The melting temperatures by Dubrovinskaia and Dubrovinsky and Goncharov et al. are somewhat lower at P>25 GPa but with the experimental uncertainties considered, these curves are fairly consistent with our estimates.

Figure 7 shows the densities of solid ice VII and fluid H2O along the melting curve together with the data of Goncharov et al. The densities of ice VII and fluid H2O are calculated from the EoS in this study and by Belonoshko and Saxena, respectively. Note that the melting temperatures of this study and Goncharov et al. are not exactly the same and that their density data are not on the equilibrium, namely, below and above the melting-T for ice VII and fluid H2O. The density of fluid H2O from Belonoshko and Saxena is consistent with that from Goncharov et al. as they discussed. Goncharov et al. estimated the density of fluid from the XRD data. This consistency in the volume of fluid H2O may justify both theoretical and experimental estimations. In contrast, the density of solid ice VII from our EoS is larger than those from Goncharov et al. Since Goncharov et al. used a laser-heating technique for high-temperature runs, the uncertainty in temperature is large in their experiments. Accordingly, calculated pressures upon heating through a pressure standard may also have a large uncertainty. However, both Goncharov et al. and this study show a similar behavior of the volume change at melting with increasing pressure (Fig. 7). Above P=10 GPa, the volume change is almost constant in both studies although Goncharov et al. claimed it was slightly increased at higher pressures. This suggests that the compressibility of fluid H2O is close to that of solid ice VII above P=10 GPa. Thus, the physical properties of two phases will be similar at high pressures. However, at higher pressures (at least P > 50 GPa from this study), the dynamically disordered phases will be stabilized suggesting that experimental constraints on the EoS for those high-pressure ices should be made in the near future.

The enthalpy change upon melting was calculated through thermodynamics as

$$\Delta H_m = T \Delta S_m,$$

where $\Delta H_m$ and $\Delta S_m$ are an enthalpy and entropy changes on melting, respectively. Figure 8 shows the calculated $\Delta H_m$ together with data from Goncharov et al. The $\Delta H_m$ in this study is somewhat smaller (60 kJ/mol at P=40 GPa) than
Therefore, the large $\Delta H_m$ in Goncharov et al.\textsuperscript{26} is not likely to be the consequence of the dissociation of H$_2$O in the fluid. Since they derived $\Delta H_m$ through the relation of $dP/dT = \Delta H_m/\Delta V$, the larger $\Delta H_m$ might come from the smaller $\Delta V$, i.e., smaller density of ice VII (Fig. 7). If the density of ice VII from our EoS is used, $\Delta V$ will be larger and accordingly, $\Delta H_m$ will be smaller. Nevertheless, the possibility of the dissociation of fluid H$_2$O still remains as Goncharov et al.\textsuperscript{26} demonstrated, independently of the $\Delta H_m$ calculations. As they discussed, if there are species of H$_3$O$^+$ and OH$^-$ in addition to H$_2$O in the fluid phase, the system is no longer unary. Then, there will be a phase loop at the melting of ice, giving rise to a complex structure of the icy planets.

In conclusion, we collected volumetric data of ice VII at simultaneous high-$P$-$T$ conditions with the angle-dispersive high-resolution XRD technique. The thermal EoS for ice VII was constructed in the Anderson-Grüneisen formulation. Thermodynamic calculation of the melting of ice VII shows that the melting line is located at temperatures higher than the previous calculation due to the smaller molar volume of ice VII in this study. Estimated $P$-$T$ range of the melting line is consistent with the recent experimental data. The $P$-$V$-$T$ data of ice VII and the melting curve estimated in this study provide key information for the discussions of the structure of icy planets, condensed matter physics, and material sciences.

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