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Intensification of the meridional temperature gradient in the Great Barrier Reef following the Last Glacial Maximum

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Tropical south-western Pacific temperatures are of vital importance to the Great Barrier Reef (GBR), but the role of sea surface temperatures (SSTs) in the growth of the GBR since the Last Glacial Maximum remains largely unknown. Here we present records of Sr/Ca and δ18O for Last Glacial Maximum and deglacial corals that show a considerably steeper meridional SST gradient than the present day in the central GBR. We find a 1–2 °C larger temperature decrease between 17° and 20°S about 20,000 to 13,000 years ago. The result is best explained by the northward expansion of cooler subtropical waters due to a weakening of the South Pacific gyre and East Australian Current. Our findings indicate that the GBR experienced substantial meridional temperature change during the last deglaciation, and serve to explain anomalous deglacial drying of northeastern Australia. Overall, the GBR developed through significant SST change and may be more resilient than previously thought.

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Sea surface temperature (SST) gradients in the south-western tropical Pacific control the extent of the Western Pacific Warm Pool (WPWP, SST >28°C), the position and intensity of the Australian summer monsoon and, consequently, hydrological changes on the adjacent continent. This region hosts the world’s largest extant coral reef, the Great Barrier Reef (GBR) World Heritage Area, with its unique ecosystem that has evolved over hundreds of thousands of years, apparently in response to major environmental perturbations. The SST rise from the Last Glacial Maximum (LGM) to the present is well constrained in the equatorial Pacific Ocean, but is relatively unknown for the unique GBR region. A previous study suggested regional SST changes of ~1.5°C or less during the late Pleistocene, questioning warming SST as an explanation for the rise of the GBR. However, this result is inconsistent with SST reconstructions for the southeastern Coral Sea and the WPWP that indicate a ~3°C cooling during the LGM. Clearly, a better understanding of the temperature history of the GBR ecosystem since the LGM is essential in order to establish a baseline against which to judge the potential response of this region to future climate change.

Here we investigate the Sr/Ca and δ18O environmental proxies in precisely U-Th dated fossil shallow-water corals drilled by Integrated Ocean Drilling Program (IODP) Expedition 325 along the shelf edge seaward of the modern GBR to reconstruct the meridional temperature gradient since the LGM (Fig. 1). Sr/Ca variations in aragonitic coral skeletons are a proxy for SST variability, and δ18O can be influenced by site-specific changes in seawater δ18O. We show that the meridional SST gradient in the GBR region was considerably steeper than today during the LGM and last deglaciation, which is best explained by northward expansion of cooler subtropical waters, owing to weakening of the South Pacific subtropical gyre and East Australian Current (EAC). Our findings indicate that the GBR experienced substantial and regionally differing temperature change during the last deglaciation, much larger temperature changes than previously recognized. Furthermore, our findings suggest a northward contraction of the WPWP during the LGM and last deglaciation, and serve to explain anomalous drying of northeastern Australia at that time.

Results
Coral preservation and ages. The drill cores of IODP Expedition 325 intersected massive and robust branching/columnar Isopora palifera/cuneata colonies, which are common in shallow-water (0–10 m), high-energy reef crest environments. Fossil Isopora corals were recovered at Noggin Pass (NOG, 17.1°S) and Hydrographer’s Passage (HYD, 19.7°S) in the central GBR (Fig. 1). The corals were screened for possible diagenetic alteration of their skeletons using X-radiography, powder X-ray diffraction, thin-section petrography and Mg/Ca screening, and are well preserved (Methods and Supplementary Figs 1–4). U-Th dating yielded coral ages spanning 25 to 12 thousand years before the present (kyr BP; ‘present’ is defined as AD 1950) (Methods and Supplementary Table 1). All coral specimens were analysed for Sr/Ca, δ18O, Mg/Ca and δ13C along their major growth orientation (Methods; Supplementary Figs 1, 3 and 4; Supplementary Table 2). A total of 881 samples were analysed at subseasonal resolution in 7 fossil Isopora corals (Supplementary Fig. 3). An additional 18 fossil and 13 modern Isopora were analysed for bulk geochemical composition, as the major aim of the study is the reconstruction of mean SST changes from geochemical averages of corals.

Coral-based SST reconstruction. The performance of Sr/Ca and δ18O in Isopora corals was assessed using modern colonies from Heron Island (HER, 23.4°S) in the southern GBR (Fig. 1), which provides a modern analogue with relatively cool SSTs such as those experienced by last deglacial central GBR corals. The modern corals reveal between-colony offsets in mean Sr/Ca and δ18O that are equivalent to ~1–2°C (Fig. 2) and similar to those seen in Porites corals (Supplementary Note 1). We note that the modern Isopora corals represent a range of shallow-water habitats on the reef, comparable to the environment represented by the fossil corals. Thus, the between-colony offsets reflect the combined effects of differences in water depth, location on the reef and different time intervals. Initial comparison of bulk samples from modern Isopora corals from Papua New Guinea, and the central and southern GBR
Figure 2 | Great Barrier Reef coral Sr/Ca and δ18O and western tropical Pacific temperatures. (a) Western Pacific Warm Pool (WPWP) sea surface temperature (SST) anomaly (left y axis) and southeastern Coral Sea SST10 (right y axis) reconstructed from planktonic foraminiferal Mg/Ca. Modern annual mean WPWP SST is >4 °C warmer than SST in the southeastern Coral Sea35. (b) Mean Sr/Ca of individual Isopora palifera/cuneata corals from Noggin Pass (NOG) and Hydrographer’s Passage (HYD), central Great Barrier Reef (GBR), and Heron Island (HER), southern GBR (approximate latitude indicated). Weighted least-squares regression lines are shown for NOG and HYD, utilizing data variances as weights. HYD coral Sr/Ca is significantly different from NOG coral Sr/Ca (Methods and Supplementary Fig. S5). The coral-based SST anomalies are not adjusted for changes in seawater Sr/Ca, and thus provide upper estimates of the magnitude of cooling. However, the effects of seawater Sr/Ca changes are similar at both sites, so the reconstructed SST differences between sites are not affected. For reference, the fossil coral-based Sr/Ca-SST anomalies are plotted relative to average Sr/Ca at HER (dashed green line). Modern mean SST35 at NOG (26.6 °C) and HYD (26.0 °C) is shown relative to SST at HER (24.5 °C) and scaled using the mean coral Sr/Ca-SST relationships of −0.084 mmol mol⁻¹ per °C (ref. 34) (solid red and blue lines) and −0.140 mmol mol⁻¹ per °C (ref. 32) (dashed red and blue lines). (c) As in b, but for mean coral δ18O corrected for the influence of changes in ice volume using a global compilation of benthic foraminifer δ18O records37 (for uncorrected coral δ18O see Supplementary Fig. 8). The resulting coral δ18O-SST anomalies are shown relative to average δ18O at HER (dashed green line) using the average of three mean coral δ18O-SST relationships (−0.22% per °C, refs 32-34). The larger cooling inferred from coral δ18O compared with Sr/Ca suggests a positive seawater δ18O anomaly. Modern mean SST35 at NOG and HYD is shown relative to HER (solid red and blue lines). The grey shading indicates the timing of the Last Glacial Maximum63,64 (LGM).
seawater Sr/Ca (ref. 15). Interestingly, if coral δ18O is used on its own to estimate glacial SSTs (correcting for glacial-interglacial changes in seawater δ18O (ref. 37) but not for regional seawater δ18O changes), we find LGM SSTs ~ 7 °C cooler than today at NOG (17°S) and ~9 °C cooler during the late LGM (~19 kyr BP) at HYD (20°S). Taken together, the apparent 2–4°C difference between the coral Sr/Ca- and ecologically-based estimate (4–5°C) and δ18O-based estimate (7–9°C) of LGM SST cooling suggests a positive δ18O anomaly (0.4–0.8‰) for GBR surface waters. This anomaly would be consistent with reduced regional precipitation minus evaporation balance at the time, although other changes in the hydrological cycle and surface ocean circulation may also contribute.

Coral-based SST gradient reconstruction. Our most striking finding is that the mean Sr/Ca values for the 24 fossil *Isopora* corals with ages spanning ~20 to ~13 kyr BP for the southern (HYD) and northern sites (NOG) form two separate groups (Fig. 2; Supplementary Table 2). Regression analysis indicates that the HYD coral Sr/Ca is significantly higher than the NOG coral Sr/Ca (Methods and Supplementary Fig. 5). In addition, the significance of the difference between the HYD and NOG palaeo-SSTs is verified by regression analysis of the coral δ18O values for the two sites (Fig. 2). Importantly, our analysis of relative changes in the meridional SST gradient is not biased by glacial-interglacial changes in seawater Sr/Ca and δ18O because the effects would be essentially the same at both sites. The difference in mean SST changes given by the HYD and NOG coral Sr/Ca and δ18O indicates that the southern site was on average 2–3°C cooler relative to the northern site from ~20 to ~13 kyr BP (Fig. 2). Given the relatively small meridional difference in SST today (<0.6°C; ref. 35; Supplementary Discussion) between the two sites (Fig. 1), our results indicate a steeper meridional SST gradient along the central GBR between 17° and 20°S throughout most of the LGM and deglaciation.

Discussion

Today, meridional SST gradients such as those along the GBR during the LGM and deglaciation are only observed south of ~30°S on the eastern Australian margin (Fig. 1). At this latitude, the EAC, which transports warm tropical waters southward along the eastern Australian coast, separates from the continent to form the Tasman Front38,39. Therefore, the steeper LGM and deglacial SST gradient may be due to EAC weakening and northward expansion of cooler subtropical waters. Furthermore, the relatively high δ18O of central GBR waters during the LGM, shown by the coral records, supports the interpretation of northward expansion of high-salinity subtropical waters (Supplementary Fig. 6). The results are consistent with studies of planktonic foraminifer assemblages40 and δ18O (ref. 38), indicating northward displacement of subtropical waters, the EAC separation and the Tasman Front to ~25°S during the LGM40 and deglaciation38. However, our results indicate that cooler waters were displaced as far north as 20° to 17°S in the central GBR, and that the steeper meridional SST gradient at this latitude was a robust feature of the LGM and deglaciation. The coral Sr/Ca data indicate that the SST gradient decreased significantly from the LGM towards the late deglaciation (Fig. 2; Supplementary Fig. 5), due to greater warming at the southern site compared with the northern site.

The available palaeoclimate records indicate that the steeper meridional SST gradient in the central GBR does not reflect a regionally steeper meridional SST gradient in the south-western tropical Pacific because the south-western tropical Pacific meridional SST gradient was actually weaker than today during the last deglaciation. The deglacial SST rise in the southeastern Coral Sea (MD97-2125; 161°44'E, 22°34'S) started earlier (~20 kyr BP) and reached modern SST as early as ~17 kyr BP (ref. 10), whereas warming in the WPWP commenced at ~17.5 kyr BP and reached modern SST as late as ~12 kyr BP (ref. 7) (Fig. 2). Both regions cooled by similar amounts during the LGM (~3°C) (refs 7,10) and today the WPWP is >4°C warmer than the southeastern Coral Sea35 (Figs 1 and 2), so the earlier start of the deglacial warming in the southeastern Coral Sea, relative to warming in the WPWP, means that the meridional SST gradient in the tropical southwest Pacific was weaker than today during the last deglaciation. This is opposite to our reconstructed steepening of the meridional SST gradient along the central GBR at that time, and is consistent with weakening of the EAC along the eastern Australian coast during the LGM and deglaciation. Furthermore, the reduction in heat transfer from low to mid-latitudes due to weakening of the EAC, which is the western boundary current of the South Pacific subtropical gyre, may have played an important role in dampening the amplitude of LGM cooling in the WPWP. Interestingly, the increase in heat transfer from low to mid-latitudes implied by EAC strengthening after ~12.7 kyr BP, suggested by the coral records, coincides with the end of the deglacial SST rise in the WPWP. The results suggest an important role for variations in the strength of the South Pacific's western boundary current, and changes in the transport of warm tropical waters to higher latitudes, in contributing to SST changes in the WPWP on glacial-interglacial timescales.

A steeper meridional SST gradient and cooler ocean surface off northeast Australia between ~20 and ~13 kyr BP may explain anomalously dry conditions in northeastern Australia at that time inferred from the Lynch's Crater sediment record41 (145°70'E, 17°37'S). The cooler deglacial waters could have reduced the effectiveness of the southeasterlies in advecting moisture from the Coral Sea to the adjacent continent. Furthermore, these waters could have potentially restricted the southern boundary of the WPWP and southward migration of the Intertropical Convergence Zone to a more northerly position, thus inhibiting development of the Australian summer monsoon. The southward displacement of subtropical waters after ~12.7 kyr BP, indicated by the coral records, was potentially accompanied by a southward shift of the southern boundary of the WPWP and the Intertropical Convergence Zone, and was probably driven by spin-up of the South Pacific subtropical gyre and corresponding strengthening of the EAC in response to deglacial warming.38

Our findings indicate that the GBR has experienced much larger SST changes—both spatially and temporally—since the LGM than previously recognized. This was especially the case in the south, owing to significant steepening of the meridional SST gradient resulting from a northward expansion of cooler subtropical waters during the LGM and deglaciation. Combined with evidence for the existence of extensive reefs along the shelf edge seaward of the modern GBR at that time11,42, our results provide new insights into the ability of coral reefs to adapt to temperature change. Most corals are well adapted to their local temperature regime, but are susceptible to thermal stress resulting in bleaching and sometimes mortality, where the mean summer maximum temperature is exceeded by > ~1 °C (ref. 43). As a consequence, the absolute thermal limit for coral bleaching is strongly site dependent43, and questions remain on the timescales over which corals may adapt to temperature rise44,45. The apparent resilience of the GBR *Isopora* coral reef community throughout much of the last deglaciation, despite substantial increases in SST from the LGM condition, demonstrates that adaption occurred over a few thousand years or less. This apparent resilience further suggests that, considering temperature
alone, southward expansion of this robust reef crest community could be an important response to future climate warming.

**Methods**

**Coral samples.** Fossil *Isopora palifera/cuneata* corals were recovered in 2010 by IODP Expedition 325 off the central Great Barrier Reef (GBR). Corals were recovered at depths between 156 and 126 m below modern sea level at Noggin Pass (NOG; 146.6°E, 17.1°S; IODP sites M0033 and M0057) and Hydrographer’s Passage (HYD; 150.3°E, 19.7°S; IODP sites M0031, M0033, M0035, M0036 and M0039) using the mission-specific platform ‘Greatship Maya’. Modern *Isopora palifera/cuneata* corals were collected at Heron Island (HER; 151.9°E, 23.4°S) in the southern GBR from 1964 to 1979 at depths between 0 and 14 m.

**Screening for diagenesis.** The aragonitic skeletons of all fossil corals were analysed by powder X-ray diffraction and X-radiography, indicating ≤ 1.5% calcite (Supplementary Table 2) and demonstrating that skeletal areas of obvious diagenetic alteration were not analysed (Supplementary Fig. 1). Selected fossil corals were also investigated using petrographic thin sections and showed excellent preservation of primary porosity, with no evidence for significant amounts of secondary aragonite or calcite cements (Supplementary Fig. 2). Skeletal Mg/Ca ratios indicate the absence of significant amounts of high-Mg calcite and secondary aragonite cements along the analysed transects (Supplementary Figs 3 and 4). Seven corals were analysed at 0.3-mm resolution for element/Ca and stable isotope ratios. Continued, subtle diagenetic alteration was detected on this spatial scale. The reported mean geochemical values are similar to those for corals of similar age analysed for bulk composition.

**U-Th dating.** The chronology of the fossil corals was determined by U-Th dating carried out using different methods at three laboratories: Woods Hole Oceanographic Institution (WHOI), Australian National University (ANU) and University of Oxford (OX). The originating laboratories are identified alongside the data in Supplementary Table 1. U–Th ages are reported in yr before the present relative to AD 1950. Interlaboratory replicate age determinations for a single specimen agree to within 100 years, which is similar to the intra-coral variability observed in some specimens measured using the high-precision (WHOI) method. The corals used in this study have initial 234U/238U activity ratios between 1.480 and 1.505, which are close to the modern seawater composition giving additional confidence in the chronology. Briefly the methods are:

1. **At WHOI,** U and Th isotopes are measured by MC-ICP-MS in static mode with all isotopes in Faraday collectors. Large ~ 5-g subsamples of coral are dissolved and spiked with a mixed 233U:234U:235U tracer, optimized for deglacial age samples and co-precipitated with Fe. To determine the 235U/238U activity, purified U and Th fractions are recombined such that U and Th are measured at isotopic ratios that can be closely matched to bracketing standards. The 235U/238U is similarly determined statically in Faraday collectors but on an unsupplied aliquot.

2. **At ANU,** U and Th isotopes are measured by thermal ionization mass spectrometry (TIMS) in static mode with 233U:234U:235U tracer, optimized for deglacial age samples and co-precipitated with Fe. To determine the 235U/238U activity, purified U and Th fractions are recombined such that U and Th are measured at isotopic ratios that can be closely matched to bracketing standards. The 235U/238U is similarly determined statically in Faraday collectors but on a 238U:235U tracer, and U and Th are separated using U-Te via a single pass. The 237-239-232Th isotopes were measured simultaneously in charge mode, in Faraday cups, with 20 pF feedback capacitors as active electrometer elements. At ANU, 233-234-235U are also measured in charge-collection mode, in Faraday collectors, using 20 pF feedback capacitors as active electrometer elements. The instrument was calibrated with reference to a 232Th-U beam intensity was kept between 8 × 106 and 5 × 107 ions per second. At ANU, 233-234-235U isotopes are measured separately: U isotopes, statically; and Th by peak hopping the 229 and 232 peaks, independently, in Faraday collectors, using 20 pF feedback capacitors as active electrometer elements. Extensive measurements with an unsupplied U-standard HU-1 showed that the curvature of the 238U tail remained invariant under a range of operating conditions, particularly when the 233-235U peaks were dominant. U sample loads on single rhinestones, filamented, from 0.5 to 0.8 µg and the 238U beam intensity was kept between 8 × 1011 and 10 × 1011 for A hours. At these intensities, the 205Th–231mTh beam was used to avoid response-time problems encountered with the considerably slower 205Th beam. The instrument response was calibrated with reference to a secular equilibrium standard HU-1. Comparisons with Western Australian coral samples for the last interglacial 47 showed precise agreement with previous measurements.

3. **At OX,** U and Th isotopes are measured by MC-ICP-MS utilizing ion counter collectors for the minor isotope beams. Approximately 0.3 g of coral samples are dissolved and spiked with a mixed 236U:238U tracer. U and Th are purified and measured separately: U isotopes, statically; and Th by peak hopping the 229 and 230 beams into an ion counter normalizing beam intensity between steps with either 229Th or 231Th U–Th measured in Faraday collectors. Instrumental biases and relative collector efficiencies are accounted for using standard-sample bracketing using U and Th iso  

deposit is calculated relative to AD 1950 for consistency.

**Microsampling.** All coral specimens were microsampled continuously along their major growth orientation (Supplementary Fig. 1). We note that the major growth axis of a coral colony is difficult to discern from drill cores with a diameter much smaller than the colony, and that sampling paths with respect to coral morphology may produce different geochemical variations. We cannot completely exclude such effects but note that the majority of our massive *Isopora* specimens indicate uniform parallel growth of corallites, suggesting the absence of colony morphology effects on this spatial scale. Seven corals were microsampled at a subseasonal resolution, by milling a trench 2.4 mm wide and 3.0 mm deep, using a sample-step increment of 0.3 mm. All other corals were sampled for bulk geochemical composition by milling a continuous trench of bulk width and depth.

**Geochemical and isotopic analyses.** Coral element/Ca and stable isotope ratios were measured on splits of the microsampled coral as an SST archive from HYD and NOG and the difference in straight-line fit resulted. The s.e. of these differences (versus time) was used to generate the error envelope (Supplementary Fig. 5). Simulations (400) were performed with independent, normally distributed random deviations (s.d. equal to dating error for sampled corals) indicate that each coral analysed encompasses several years of skeletal growth. On the basis of this analysis, there is no clear indication for systematically lower growth rates at the southern site compared with the northern site. Thus the relatively cool reconstructed SSTs at the southern site are not the result of a potential influence of lower growth rates on SST proxies.

**Isopora coral as an SST archive.** *Isopora* has been considered a subgenus of the coral genus *Acropora,* but was recently elevated to genus57. Massive and robust branching *Isopora palifera/cuneata* corals are recovered at depths between 10 m of well-flushed reef crest environments23,24, providing an excellent potential archive of open-ocean SST. However, submillimetre-scale analysis of seven massive *Isopora,* the growth form most common in our study, did not reveal clear annual cycles in Sr/Ca and δ18O (Supplementary Fig. 3). It is possible that the complex skeletal architecture, including discontinuous filament-like structures perpendicular to the growth direction, and spatially heterogeneous calcification processes within massive *Isopora* corals obscure the annual cycle that diagnosis was ruled out for these samples. We note that studies on branching *Acropora* corals (A. palmata, A. nobilis) did not reveal clear annual cycles in Sr/Ca and δ18O along axes of radial growth58,59, which was attributed to spatially heterogeneous calcification processes such as secondary thickening of skeletal elements60–62.

In the absence of definitive calibration data for massive *Isopora,* the average of well-established mean proxy-SST relationships for *Porites* corals was used to reconstruct changes in mean SST. For Sr/Ca, on calibrations derived from bulk coral values for reef settings with different average temperatures33 and on annual mean calibrations using coral time series52. Preliminary bulk calibration experiments on modern *Isopora* corals suggest similar *Porites* and 8°-O-SST relationships to the better-studied *Porites* genera (Supplementary Fig. 7; Supplementary Note 2). The *Isopora* bulk calibration slope for Sr/Ca lies closer to the seasonal *Porites* slope for Sr/Ca than to the bulk *Porites* slope. Furthermore, seasonal Sr/Ca calibration slopes for *Porites* were ANU (Canberra), LDEO (Palisades), Geosciences (Edinburgh), AIST (Tsukuba) and AORI (Kashiwa), using ICP-OES, ICP-MS and dual-inlet stable-isotope ratio mass spectrometry. Bulk powders of fossil corals were analysed at least in triplicate. The coral reference material JC1-1 was used to check for interlaboratory offsets64. The JC1-1 concordance value used in this study is 0.71% (ref. 34, 55) for Sr/Ca and 4.75% for δ18O, 2.522% for Mg/Ca and 1.58% for δ54Ca. These values represent the average of > 100 element/Ca and > 60 δ18O and δ54Ca analyses of splits of JC1-1 powder that were treated like samples in the participating laboratories.

**Statistical analysis.** Regression analysis was performed by applying a straight-line weighted least-squares fit (utilizing data variances as weights, 1/σ(δ)², ref. 61) to the coral data from HYD and NOG and the difference in straight-line fit (HYD – NOG) was determined. A Monte Carlo error simulation65 was performed for the difference in straight-line fit by generating random deviations from original data points, from which simulated straight lines and simulated fit differences resulted. The s.e. of these differences (versus time) was used to generate the error envelope (Supplementary Fig. 5). Simulations (400) were performed with independent, normally distributed random deviations (s.d. equal to dating error for age simulation and equal to measurement error × inflation factor for proxy simulation). The fit residuals were used to determine the mean deviation from the fitted line (root of residual mean square). This quantity was compared with the mean proxy measurement error, which is smaller by a factor of 2 for HYD coral
$\delta^{18}O$, smaller by a factor of 3 for NOG coral $\delta^{18}O$, smaller by a factor of 2 for HYD coral Sr/Ca and about the same for NOG coral Sr/Ca. Consequently, a conservative approach was chosen by artificially inflating the assumed measurement error in order to simulate the influence of errors induced by effects such as between-colony variability. A measurement-error inflation factor of 3.0 was used to calculate the s.e. band for our difference in straight-line fit. However, the difference is still larger than zero for both coral Sr/Ca and $\delta^{18}O$. Therefore the HYD coral Sr/Ca is significantly different from the NOG coral Sr/Ca, and the HYD coral $\delta^{18}O$ is significantly different from the NOG coral $\delta^{18}O$.

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Author contributions
T.F., H.V.M., B.K.L., A.W.T., A.S., M.K.G., M.T. and Y.Y. were responsible for coral geochemical analysis; A.L.T., T.M.E. and W.G.T. carried out coral U-Th dating; D.C.P. provided modern corals; M.M. performed statistical analyses; Y.Y. and J.M.W. were Expedition 325 co-chief scientists; T.F., B.K.L., A.W.T., M.K.G., A.S., M.I., A.L.T., T.M.E., W.G.T., M.T. and D.C.P. were Expedition 325 scientists; T.F. was responsible for data compilation and wrote the manuscript together with H.V.M., M.K.G., A.W.T. and B.K.L.; all authors contributed to data interpretation and manuscript preparation.

Additional information
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